



Figure 7. Derived and measured electric field power spectra. (a) from balloon, $L > 6$ [Mozer, 1971]; (b) from balloon, $L > 6$ [Holzworth and Mozer, 1979]; (c) from whistler duct convection, each curve represents one duct, $L = 2.3$ [Andrews, 1980]; (d) from radar, $L = 1$ [Earle and Kelley, 1987]; (e) from satellite, $L = 6.6$ [Junginger et al., 1984]; (f) present results.

currents, and therefore affects the high-latitude convection [Senior and Blanc, 1984], which may, perhaps through a neutral wind scenario like that of Blanc and Richmond [1980] and Blanc [1983] be responsible for the day-to-day changes in the low-latitude dynamo. In addition, as the distance between the effective inner edge of the ring current and the ionosphere decreases, the diamagnetic effect of the ring current magnetic field increases, as well as the energy input in the form of precipitating energetic neutrals [Tinsley, 1979].

If any of these (admittedly small) inputs modify the ionospheric electric fields sufficiently, then this feedback mechanism may account for some of the fluctuation power of the dynamo fields. One possible scenario would begin with a high-latitude electric field penetrating to low latitudes and enhancing the azimuthal electric field either globally or locally. This static field would cause the ring current to move inward, enhancing the diamagnetic effect as well as the precipitating neutral flux. Both effects would increase the ionospheric conductivities, which in turn reduce the dynamo electric field strength. The net result is that a pulsation of the electric field is generated, increasing the internal contribution of the electric diffusion coefficient, and intensifying the inner edge of the ring current. Thus the next electric field pulsation would create an even greater response. Future research might be aimed at analyzing such processes.

possibly contribute greatly to the transport of the quiet ring current during the solar minimum period of 1985–1989. We have concentrated on data in the ~ 30 – 300 keV energy range, where diffusion, not convection, is the dominant transport mechanism. We have used a semi-empirical derivation of the presumed ionospheric origin electric diffusion coefficient to fit the data using a maximum likelihood method. We also found that above $L \sim 6$, externally driven field fluctuations dominate the diffusive transport, but below this shell, internal origin electric field fluctuations cause most of the cross- L transport. Additionally the deduced radial gradient of the “internal” fluctuations suggest that either the field fluctuation source itself has strong latitudinal gradients and/or that the surmised ionospheric fields appear to be shielded from the magnetosphere. Assuming that magnetic field lines are also electric field equipotentials and assuming that diffusion is caused by only the lowest spatial harmonic, we can estimate the power level of these fluctuations in the ionosphere. We find that this estimate is comparable in magnitude, though not fully consistent in spectral form with the few ionospheric measurements that have been made. We speculate that the effects of the ring current–equatorial electrojet may produce a feedback (or regulation) mechanism that at least in part determine the global equilibrium of the ionospheric current system.

Acknowledgments. We thank G. Gloeckler and the AMPTE/CHEM team for much support. Special thanks to Hamilton, S. Christon, and M. Greenspan for valuable discussions and criticism. We thank S. Nylund and the AMPTE/CCE Science Data Center for their invaluable assistance. We also appreciate the careful reading of the manuscript by the referees. This work was supported by NASA grant NAG5-716, and the Swiss National Science Foundation.

The Editor thanks J. U. Kozyra and another referee for their assistance in evaluating this paper.

References

- Anderson, P. C., R. A. Heelis, and W. B. Hanson, The ionospheric signatures of rapid subauroral ion drifts, *J. Geophys. Res.*, **96**, 5785–5792, 1991.
- Andrews, M. K., Power density of equatorial electric field at $L = 2.3$, *J. Geophys. Res.*, **85**, 1687–1694, 1980.
- Baumjohann, W., G. Haerendel, and F. Melzner, Magnetospheric convection observed between 0600 and 2100 LT: Variations with K_p , *J. Geophys. Res.*, **90**, 393–398, 1985.
- Bevington, P. R., *Data Reduction and Error Analysis for the Physical Sciences*, McGraw-Hill, New York, 1969.
- Blanc, M., Magnetospheric convection effects at mid-latitudes: Theoretical derivation of the disturbance convection pattern in the plasmasphere, *J. Geophys. Res.*, **88**, 235–251, 1983.
- Blanc, M., and A. D. Richmond, The ionospheric disturbance dynamo, *J. Geophys. Res.*, **85**, 1669–1686, 1980.
- Carpenter, D. L., New whistler evidence of a dynamo origin of electric fields in the quiet plasmasphere, *J. Geophys. Res.*, **82**, 1558–1564, 1978.
- Carpenter, D. L., and R. R. Anderson, An ISEE/whistler model of equatorial electron density in the magnetosphere, *J. Geophys. Res.*, **97**, 1097–1108, 1992.